# Improving WRF simulations of coastal storms with better water vapour initial fields from InSAR interferometry

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# Outline

Coastal storms can be extreme, due to high input of water vapor and strong mesoscale forcing (topography)

A state of the art model, such as WRF, such be able to deal with those systems (at the right resolution)

However one needs a good initial state, and the water vapor field may be critical. Here:

We estimate PWV maps by Sentinel-1 data (and GNSS)

> WRF 3dVar assimilation experiments

### GNSS, InSAR and water vapor

GNSS (GPS) data is a new source of water vapor data at relatively coarse resolution (50 km) but with frequent sampling (minutes to 1 h)

We propose to use an even newer source of data, InSAR (Synthetic Aperture Radar) interferometric maps, produced by Sentinel 1a,b

2-SAR images are available (now) every 6 days (ascending+descending)

InSAR images (made from 2 co-located SAR maps) are produced to infer land movements, but are affected by water vapor (as GNSS): if there are no land movements we may infer anomalies of water vapor.

### The Adra Storm (Sep 2015)

Two severe weather events over Adra (Almeria, Spain): 6<sup>th</sup>-7<sup>th</sup> September 2015 <u>with 12-13h apart</u>

50.0°1

45.0°N

40.0°N

6<sup>th</sup> Sep. 2015 - 22h 7<sup>th</sup> Sep. 2015 - 11h 35.0°N RADAR RADAR WRF Model WRF Model ADRA ADRA 2°W 2°W not predicted by NWP models, each ~6h after Sentinel!



# The Adra storm



# Questions

- What impact would GNSS data at the available resolution would have in the forecast?
- Is there any added value in the much higher resolution InSAR data?
- What dynamical changes are produced by the new data: if we get a storm (we do!) why is that?

# The SAR product



Synthetic Aperture Radar  $s_1 = A.e^{(j\phi_B)}.e^{(-j(\frac{4\pi}{\lambda}).r_1)}$ 

# InSAR: SAR interferometry

2 SAR images with the "same" view (now 6 days apart)

Interferogram

$$s = s_1 s_2^* = |s_1| |s_2| e^{-j \left(\frac{4\pi}{\lambda}\right) \cdot R_1 + j \left(\frac{4\pi}{\lambda}\right) \cdot R_2}$$

$$s = s_1 s_2^* = |s_1| |s_2| e^{-j \left(\frac{4\pi}{\lambda}\right) \cdot R_1 + j \left(\frac{4\pi}{\lambda}\right) \cdot R_2}$$
Phase difference
$$\phi_I = -\frac{4\pi}{\lambda} (R_1 - R_2)$$
If the land is quiet it is a

difference in water vapor (PWV)

### Previous work @IDL: ΔPWV maps from InSAR

If the terrain deformations can be neglected, hydrostatic and ionospheric contributions removed, InSAR provides maps of differential slant wet delay (ΔSWD)

$$\Delta SWD \ (t_M, t_S) = \frac{\lambda}{4\pi} \Delta \varphi_{wet}$$

$$\Delta PWV \ (t_M, t_S) = \prod \Delta SWD \cdot M \left( \mathcal{G}_{look} \right) -$$

$$\Pi = 10^{-6} \cdot \rho_{H_2O} \cdot R_v \cdot \left[\frac{k_3}{T_m} + k_2'\right]$$

 $T_m$ : change with mean temperature profiles

$$T_m = \frac{\int \frac{e}{T} dh}{\int \frac{e}{T^2} dh}$$

For more details see:

P. Mateus, G. Nico, J. Catalão, **"Maps of PWV Temporal Changes by SAR** Interferometry: A Study on the Properties of Atmosphere's Temperature Profiles", IEEE Geoscience and Remote Sensing Letters, 11(12), 2065–2069, 2014.



# Previous: PWV maps by Sentinel-1 data

✓ Data with a 250 km swath at 5 m by 20 m spatial resolution (single look)
 ✓ 6 days revisiting time



### Merging and calibration are made with a small GNSS network

For more details see: P. Mateus, J. Catalão and G. Nico, "Sentinel-1 Interferometric SAR Mapping of Precipitable Water Vapor Over a Country-Spanning Area" in *IEEE Transactions on Geoscience and Remote Sensing*, vol. 55, no. 5, pp. 2993-2999, May 2017. doi: 10.1109/TGRS.2017.2658342

# Previous: PWV maps by Sentinel-1 data

### From ΔPWV to PWV



#### PWV at SLAVE date



#### PWV at MASTER date

### $PWV(master) = \Delta PWV + PWV(slave)$

### ADRA 3dVar assimilation experiments



- 1. GNSS: assimilation time at 18h and 06h (130 GNSS-PWV values)
- 3. InSAR: assimilation time at 18h (ascending orbit, 1st segment) and 06h (descending orbit, 2nd segment)

# Setup of the Adra 3dVar assimilation experiments



130 local GNSS stations used for the GNSS experiment

35 GNSS stations belonging to EUREF Network 🗲 VALIDATION

wrf3DVAR T,q

@ level 4 (~400m)



**Figure 2.** Anomalies of the water vapor mixing ratio and potential temperature at the data assimilation time

### Anomalies at assimilation time(s)



# CAPE at assimilation time(s)





Figure 4. Convective available potential energy (CAPE, in J/kg, color shading), and streamlines of the low level flow (at the  $4^{th}$  model level), before the onset of each storm: (a,d) CTRL, (b,e) A-GNSS, (c,f)

8 A-INSAR.

### The storm with InSAR assimilation (20 min)



### Validation with independent GNSS data

Statistical analysis over the 12h forecast using an independent set of **35 GNSS stations** during the 12h forecasting



### Validation against udometer

### Precipitation level measured by gauge data



- ✓ The INSAR experiments correctly forecast both, the time and intensity of rainfall. The GNSS did not significantly modified the CRTL forecast.
- ✓ Slow decay indicating an inertial effect of the WRF system when modeling quick effects as local severe precipitations having a short duration (about 3 hours).

### Control GNSS InSAR Radar



### Discussion

### <u>3dVar assimilation of InSAR-PWV</u> fields in a state-of-the-art model (WRF) can <u>improve</u>

the forecast of severe events

Improving the:
✓ Temporal scale
✓ The location
✓ Amount of precipitation

A dense GNSS network does not capture the **fine spatial details of InSAR-PWV fields**, in conditions that are favorable to the onset of deep convection

InSAR information was found to be useful for about **<u>8 hours</u>** into the simulation

Changes appear to affect the initial state of humidity, temperature and wind

However: we still only have these data every 6 days...